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Estimation of earthquake frequency content and its effect on dynamic analysis using continuous and discrete wavelet transform

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KEYWORDS

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Abstract. Wavelet transform is one of the mathematical concepts for studying the frequency content of waves. It can be divided into two groups, continuous and discrete. In general, the continuous wavelet transform is used to examine the time-frequency relationship, whereas the discrete wavelet transform is used for filtering and noise reduction in waves. In this paper, for the first time, the combination of these two concepts is used for the earthquake acceleration wave. For this purpose, eight earthquakes from four different locations in the world have been selected. Initially, each earthquake is filtered up to 5 stages using a discrete wavelet transform. At each stage of the filter, two waves of approximations and details are obtained. Due to the close approximation of the frequency content of the wave to the original earthquake, the approximate wave is used for subsequent calculations. The Fourier spectrum and the diagram of five of the dominant frequency of the earthquake are plotted in the next step. Also, using the continuous wavelet transform, the time-frequency curves of the main earthquakes and the time-frequency curves of the wave obtained from the discrete wavelet transform are investigated. The goal was to find the best stage of a discrete wavelet filter based on frequency content to reduce computations by more than 80%. The time of the strong ground motion, the structural response of a single degree of freedom, and the dynamical response of the timing of the structure of a degree of freedom are all investigated in the following step. By examining the above parameters, the best-performing wavelet transformation step is inferred.

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1. Introduction

One of the important issues in earthquake engineering is the correction of waves recorded by acceleration [1]. The recorded acceleration signals are noise-free for a number of reasons. However, the noise cannot be

extracted from these signals; therefore, the recorded earthquake wave must be corrected. Currently, the signal is corrected in a conventional way by applying the filter. The conventional and common method for making such an amendment is to use upstream and downstream filters. Earthquake recording is mostly done using seismographs and accelerators [2–5]. Seismographs are very sensitive devices that can record extremely poor earth movements caused by earthquake events with high precision. Data from seismic devices cannot be directly used in earthquake engineering, since they cannot record strong earthquakes in close proximity to earthquake centers, because at

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near distances to earthquake centers, these devices are often saturated and cannot be used for optimal recording of strong ground movements. Parameters reported by seismic stations usually include spatial and temporal coordinates, magnitude, and depth of the earthquake [6–10].

Scaling of acceleration is an important topic in earthquake engineering and structural engineering. The common goal of all studies in the field of scaling down is to reduce the dispersion of the results of dynamic analysis of the timeliness and reliability of the instrumental responses. The selected accelerator map is scaled to fit the desired design. The first studies in this field were used to scale up the maximum acceleration, speed, and displacement in acceleration maps. There are several methods for scaling the earthquake acceleration, one of them is the use of peak ground acceleration [11]. In this paper, the Maximum Earthquake Acceleration method is used to scale up the waves from the discrete and continuous wavelet transform.

Generally, earthquake waves form a set of simpler waves with different frequencies and amplitudes. The frequency of an earthquake indicates that the amplitude of these waves has expanded at different frequencies [12–15]. One of the methods for examining the frequency content is the Fourier spectrum. In this paper, the Fourier spectrum was used to study the frequency content of different waves.

The spectral analysis is performed using a curve of the structure response spectrum with a single degree of freedom. A seismic response curve, also known as a seismic response spectrum, is a curve with two axes, one horizontal and one vertical, one representing the structure's period or frequency and the other representing acceleration, pseudo-acceleration, velocity, pseudo-velocity, or displacement [16–22]. The spectral curve of a specific earthquake is derived from the analysis of structures of a single degree of freedom with different natural frequencies. Therefore, the earthquake spectral curve is one of the parameters that is affected by the frequency of the earthquake. Thus, the displacement response spectrum of the wavelet filter is compared to the displacement response spectrum of the earthquake in this research to assess the competency of the suggested approach.

An earthquake's predominant energy includes the duration of a strong ground motion. Several methods have been proposed to calculate the strong ground duration of acceleration. The total amount of time between the acceleration of the Arias, which is usually between 5 and 95% of the Arias intensity, is a significant time. In this method, Arias intensity is measured in percentages, and time is measured in seconds. An effective period is defined as the duration of an earthquake where more than 90% of the total

seismic energy exists. This time is in the range of 5% to 95% of the Arias scale of the earthquake. The uniform duration is the sum of the times when the earth acceleration exceeds a given threshold, which is commonly 5% of the maximum earthquake acceleration. Selective time is the interval between the first and last time the acceleration of the earth's motion exceeds a given threshold. The maximum acceleration value is usually 5% [23,24]. In this study, a long period of time is utilized to examine strong ground motion, which is a frequent practice among engineers.

From a mathematical point of view, the conversion into a wave provides more information about the crude wave. The earthquake wave is an unstable wave because its frequencies do not occur throughout the wave's duration and vary at different times. In their study, researchers employed the wavelet transform as a time-frequency display to determine modal parameters such as natural frequencies, damping ratios, and mold shapes of a vibrating system in the dynamics of structures and earthquake engineering. They investigated the precision of their method using a numerical example to measure the vibration of a tower by wind [25]. Then, using a discrete wavelet transform, they proposed a method to improve earthquake movements. The results showed that earthquakes with a similar response spectrum can have different energy and damage potential [26]. Wavelet transform was also used for earthquake movement analysis, and the correlations between wavelet coefficients and input energy were determined by wavelet analysis. Using these principles, they examined the time-frequency features of the movements [27].

Then, the strong ground motion parameters were calculated using wavelet transform and earthquake acceleration analysis. The results showed that with the use of a wavelet transform, the parameters of the powerful earth movement could be estimated with a negligible error [28].

Some studies examined structural damage. Damage to structures can be investigated using wavelet transform. The results of various studies show the efficiency of the wavelet in this field [29,30].

Using discrete wavelet theory, Heidari and Majidi reduced the computational effort for seismic displacement and seismic velocity curves by more than 93%. They showed that by using the Haar wavelet function, they were able to reduce the calculation error to less than 5% [31]. Using discrete wavelet theory, researchers were able to analyze the association between earthquake frequencies and the natural frequency of structures in another study. In this study, they used a wavelet transform method as an alternative to modal analysis [32].

In this research, for the first time, discrete and continuous wavelet transforms are used simultaneously

to study the frequency content of earthquakes. In time history analysis, spectrum analysis, frequency content, time-frequency analysis, and a strong ground motion, a technique to reduce calculations by over 80% is described in this study. For this purpose, wavelet transforms, vibration equations, response spectrum, February spectrum, and strong ground motion time are used. To do so, each wave of various earthquakes must first be filtered using a discrete wavelet transform with up to five stages. At each stage of the filter, two waves of approximations and details are obtained. Given that the wave of details contains low-energy waves, it can be said that the frequency content of the approximate wave is closer to the original earthquake. This is why the approximate wave is used as the main earthquake substitute. By using this method in each stage of wavelet decomposition, the number of earthquake records is reduced to half of the previous stage. As a result, we can say that the number of computations and the time it takes to complete them has been cut in half. In the next step, the acceleration map, the February spectrum, the frequency content, the dynamic response of the structure of a degree of freedom, the time-frequency diagram, the response spectrum, and the strong ground motion of the various waves are investigated. The purpose of these studies is to find the best wavelet filter stage for optimizing computations.

2. Wavelet transform

There was a wavelet transform of mother functions and a different scale [33–39]. The mother functions were used in the wavelet continuity conversion and the scale functions were used in the discrete wavelet transform to separate the frequencies. ψ and φ denote the mother functions and wavelet scale, respectively. In this paper, the wavelet transform is used discretely and continuously. Therefore, both mathematical principles are discussed below. A wavelet is an alternate, real or imaginary function with a mean of zero and a finite length. Its space is defined as $\psi(t) \in L^2(R)$. The function $\psi(t)$ is called the mother wavelet, and $L^2(R)$ represents the measurable Hilbert space of integrable quadrilateral functions. The function $\psi(t)$ in the domain and space-frequency is used to create a family of wavelets and is shown as follows:

$$\psi_{u,s}(x) = \frac{1}{\sqrt{s}} \psi\left(\frac{x-u}{s}\right). \quad (1)$$

In this case, real numbers s and u are scale parameters and wavelet transforms. For a signal whose target is a wavelet filter, the coordinate space $f(t) \in L^2(R)$ is considered, in which t specifies the spatial coordinates. The continuous wavelet transform is defined as the

inner product between the signal and the wavelet function:

$$Wf(u, s) = f, \quad \psi_{u,s} = \frac{1}{\sqrt{s}} \int_{-\infty}^{+\infty} f(t) \psi\left(\frac{t-u}{s}\right) dt. \quad (2)$$

In the above equation, $Wf(u, s)$ is called the wavelet coefficient for $\psi_{u,s}(t)$. The integral of the wavelet transform of the previous equation can be written as the following integral:

$$\begin{aligned} Wf(u, s) &= \frac{1}{\sqrt{s}} \int_{-\infty}^{+\infty} f(t) \psi\left(\frac{-(u-t)}{s}\right) dx \\ &= \frac{1}{\sqrt{s}} f * \psi\left(\frac{-u}{s}\right) = f * \bar{\psi}_s(u), \end{aligned} \quad (3)$$

where $\bar{\psi}_s(x)$ is equal to $1/\sqrt{s}\psi(x/s)$.

In this paper, the Haar mother wavelet was used. The mother function of the Haar wavelet as follows:

$$\psi(t) = \begin{cases} 1 & 0 \leq t < 0.5 \\ -1 & 0.5 \leq t < 1 \\ 0 & \text{otherwise} \end{cases} \quad (4a)$$

and its scale function is as follows:

$$\varphi(t) = \begin{cases} 1 & 0 \leq t \leq 1 \\ 0 & \text{otherwise} \end{cases} \quad (4b)$$

In the wavelet transform, there are two parameters: scale and transition. The scale parameter in the wavelet is similar to the scale in the maps. The small scale represents the wave's details, while the large scale represents the wave's lack of detail. The scale, like a mathematical function, compresses or expands a wave. The large scale corresponds to the opening of the wave and the small scale makes the wave compress. The low frequency (large scale) corresponds to the general information of the wave and the large frequency corresponds to the details of the wave information. The high frequencies for a long time were unscathed from the wave, while there were low frequencies throughout the wave. In the continuous wavelet transform, the transmission and scale parameters are continuously changed. In fact, the number of computer changes in each step is very small. This makes computing more efficient [40]. Another type of wavelet transform was called wavelet fracture transformation [41] where wavelet transformation, transmission, and scale were considered discrete. The sampling rate of the time curve points can be reduced in accordance with the Nyquist norm at high scales (low frequencies), reducing the number of computations. The Nyquist sampling rate is the minimum rate at which a continuous wave can be reconstructed. Using the lower filters, it is possible to remove the high frequencies from a wave of length (number of points) and obtain the main

wave. Low frequencies can be removed and wave details acquired using the high pass filter. The wave can be divided into two parts with high and low frequencies using the simultaneous effect of these two filters on the wave. In each step, the approaches and wave details are shown with A_j and D_j . However, the main wave is obtained from the main wave with the same number of points as the main wave, which doubles the number of wave points, at each stage, using the filters. To overcome this problem, a sample reduction is used [42]. One of the consecutive points is kept and the other is deleted when the sampling is reduced. Therefore, the number of points per wave is approximately half that of the main wave points. After this step, the number of wave points can still be reduced. Because the maximum wave energy is in its approximation and the shape of this part of the wave is more similar to the wave, the filtering operation for this part of the wave is carried out and the wave is converted into two waves. The number of points in each of these waves is about half the number of points. Although from a theoretical point of view, this can be done a number of times, it is needed in dynamic wave analyses that are very similar to the initial wave and the number of points is sufficiently small. In fact, the rapid conversion of the wavelet acts like a filter bank [42,43].

In this method, approximate waves (A_j) and detail waves (D_j) were obtained from the following relationships [42]:

$$A_j = ap_{j,k} = \sum_n s(n) h_j^*(n - 2^j k), \quad (5)$$

$$D_j = de_{j,k} = \sum_n s(n) g_j^*(n - 2^j k), \quad (6)$$

where h_j is the low pass filter and D_j is the high pass filter. The h and g filters were calculated from one step to the next using the following relationships [28]:

$$g_1(n) = g(n), \quad (7)$$

$$h_1(n) = h(n), \quad (8)$$

$$g_{j+1}(n) = \sum_k g_j(k) g(n - 2k), \quad (9)$$

$$h_{j+1}(n) = \sum_k h_j(k) g(n - 2k). \quad (10)$$

These relationships show that this method is similar to the theory of filters and that the rapid wavelet transform corresponds to the filtering of the filter bank. Its inverse is also the same as the combination of the filter bank. The filters used in the wavelet inverse transformation are shown and derived from the following relationships:

$$\tilde{h}_j(n - 2^j k) = 2^{-0.5j} \tilde{\psi}(2^{-j}(t - 2^j k)), \quad (11)$$

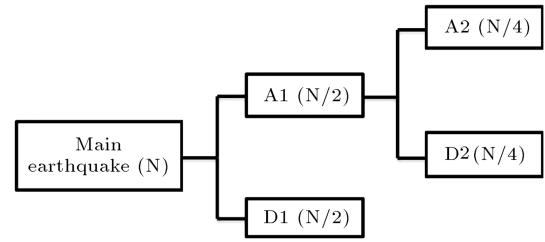


Figure 1. Down-sampling method algorithm up to two steps.

$$\tilde{g}_j(n - 2^j k) = 2^{-0.5j} \tilde{\varphi}(2^{-j}(t - 2^j k)). \quad (12)$$

The values ψ and φ are the mother and scale functions. With \tilde{h} and \tilde{g} the values and the main wave can be reconstructed using the following relationship. Figure 1 shows the wavelet transform algorithm:

$$s(n) = \sum_{j=1}^J \sum_k ap_{j,k} \tilde{h}_j(n - 2^j k) + \sum_{j=1}^J \sum_k de_{j,k} \tilde{g}_j(n - 2^j k). \quad (13)$$

3. Earthquakes used in this article

In this paper, Eight earthquakes have been used to demonstrate the applicability of the proposed technique. Table 1 shows the earthquakes used and their characteristics.

For various earthquakes, the acceleration of the original earthquake and acceleration of the approximation wave wavelet filters are shown in Figure 2. It should be noted that at each stage of the filter the number of earthquake records decreases by half. For example, the A1 to A5 records numbers in the Sarpolzahab earthquake with 19890 records are 9945, 4973, 2486, 1243, and 622, respectively.

As shown in Figure 2, the wavelet filter has good accuracy, which demonstrates the applicability of the proposed technique. With respect to Figure 2, it can be said that with increasing the filter, the approximate wave diverges from the original earthquake. For example, in Figure 2(c) and (h), the A5 wave diverges substantially from the original earthquake.

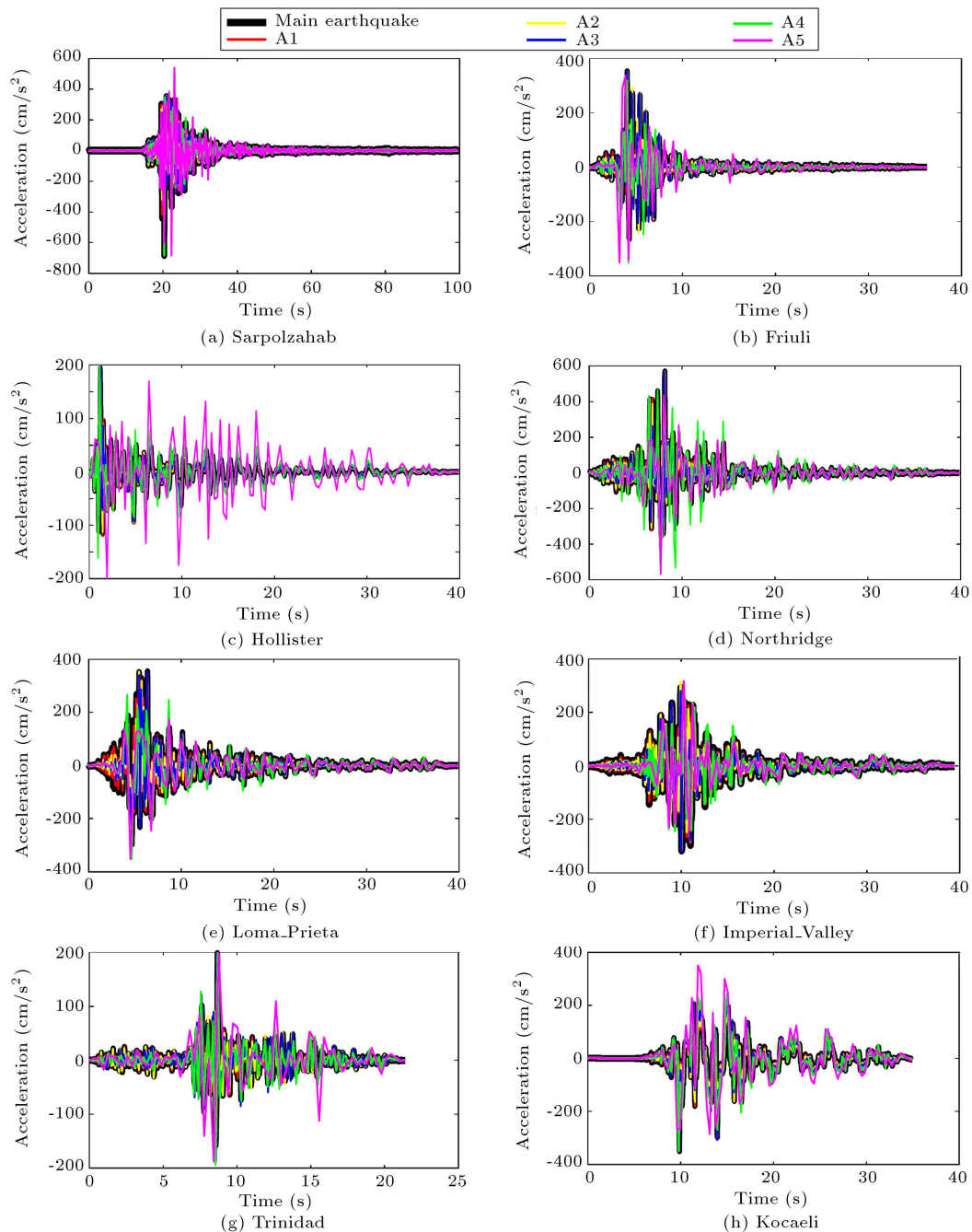
4. Fourier spectrum

Fourier Amplitude Spectrum (FAS) is function that calculate by time by using fast Fourier transform. The Fourier spectrum shows the frequency of earthquake content. The Fourier spectrum represents the dominant frequency of the earthquake. The Fourier spectrum displays the value of wave frequencies.

The desired function $S(t)$ can be shown using a set of sine and cosine functions based on the Fourier transform [44]:

Table 1. Selected characteristics of earthquake.

Earthquake	Country	Date	Earthquake magnitude	Maximum acceleration (cm/s^2)	Earthquake time (sec)	Number of records
Sarpolzehab	Iran	2017-11	7.3	684.42	99.45	19890
Friuli	Italy	1976-05	6.5	351.30	36.32	3632
Hollister	USA	1961-04	5.6	194.80	39.93	3993
Northridge	USA	1994-01	6.7	568.30	39.88	3988
Loma_Prieta	USA	1989-10	6.9	367.40	39.90	3990
Imperial_Valley	USA	1979-10	6.5	315.20	39.48	3948
Trinidad	USA	1983-08	3.2	193.60	21.40	2140
Kocaeli	Turkey	1999-08	7.6	349.00	34.96	3496

**Figure 2.** Acceleration curve of the main earthquake wave and approximate waves.

$$S(t) = a_0 + a_1 \cos 1\bar{\omega} + a_2 \cos 2\bar{\omega} + \cdots + a_n \cos n\bar{\omega} \\ + b_1 \sin 1\bar{\omega} + b_2 \sin 2\bar{\omega} + \cdots + b_n \sin n\bar{\omega}. \quad (14)$$

The previous function can be rewritten as follows:

$$S(t) = 2 \sum_{n=1}^{\infty} (a_n \cos(n\bar{\omega}t) + b_n \sin(n\bar{\omega}t)), \\ S(t_j) = \sum_{n=0}^{N-1} C_n e^{-2\pi i(nj/N)}, \\ j = 0, 1, 2, \dots, (N-1), \quad (15)$$

$$a_n = \frac{1}{T} \sum_{j=0}^{N-1} S(t_j) \cdot \cos(n\bar{\omega}t_j) \Delta t, \quad (16)$$

$$b_n = \frac{1}{T} \sum_{j=0}^{N-1} S(t_j) \cdot \sin(n\bar{\omega}t_j) \Delta t, \\ n = 0, 1, 2, \dots \quad (17)$$

$$C_n = \frac{1}{N} \sum_{j=0}^{N-1} F(t_j) e^{-2\pi i(nj/N)}, \\ N = 0, 1, 2, \dots, (N-1), \quad (18)$$

$$T = N \cdot \Delta t. \quad (19)$$

In the previous functions, N is the number of records of a discrete function (earthquake), Δt is the time step of the earthquake, T is the period of the function, and $\bar{\omega}$ is the frequency. Relation between frequency and period can be written as follows:

$$\bar{\omega} = \frac{2\pi}{T}. \quad (20)$$

$n\bar{\omega}$ can be written as follows:

$$\bar{\omega}_2 = 2\bar{\omega}, \quad \bar{\omega}_3 = 3\bar{\omega}, \dots, \quad \bar{\omega}_n = n\bar{\omega}. \quad (21)$$

FAS can be written as follows:

$$\text{FAS}(\bar{\omega}_n) = \sqrt{a_n^2 + b_n^2} = \sqrt{\text{Re}^2 + \text{Im}^2}. \quad (22)$$

In the previous functions, Re is the real part of C_n and Im is the imaginary part of C_n . In Figure 3, the Fourier spectrum of various earthquakes and wavelet filters are plotted.

As shown in Figure 3, the frequency range of the waves from the wavelet filter to the second stage is well approximated. This means that the A2 wave has a good approximation to the original earthquake after reducing 75% of the recordings.

5. Dominant frequencies of waves

In this section, five dominant frequencies in various

earthquakes and wavelet filters are investigated. The Fourier spectrum of each wave is used to develop the diagram in Figure 4. The purpose of this work is to study the frequency of waves.

Figure 4 shows the frequency number on the horizontal axis and the frequency value on the vertical axis. Figure 4 shows that the first five frequencies of the main earthquakes are similar to those of their wavelet filters. With a reduction of 75% in calculations, the A2 wave can be claimed to be almost effective in all earthquakes. Also, except for Figure 4(a) and (b), it can be said that up to the fourth stage of the filter (wave A4), with 94% reduction in calculations, is predicted with relatively good approximation of the predominant frequencies.

6. Time-frequency diagram of various earthquakes

Using the definition of discrete wavelet transform and the continuous wavelet transform, the time-frequency diagram of the waves is plotted. Its worth noting that the goal of these graphs is not to find the dominant frequencies; rather, the goal is to figure out when the dominant frequencies will appear.

A 97% reduction in well-balanced computing predicts the time of dominant of frequencies up to A5 in all earthquakes, as shown in Figure 5. The duration of a higher-energy earthquake is depicted by this curve.

7. Time of strong ground motion

In this article, the significant time of the earthquake is investigated. The use of the concept of significant time has been chosen to reduce the earthquake records and also eliminate the noise at the beginning and the end of the earthquake. In Table 2, the start and end time is shown by the strong land motion. As shown in Table 2, wavelet filters can accurately predict the start of strong ground motion in all earthquakes. Furthermore, up to the fifth step filter, the time of the end of the strong ground motion is well predicted in almost all earthquakes (A5).

Using the numbers given in Table 2, the wave associated with strong ground motion is shown in Figure 6.

Figure 6 shows that, with the exception of the A5, the technique used in all phases of the filter is consistent with the original earthquake. Therefore, the A4 wave can be used by reducing 97% of the calculations, instead of the main earthquake.

8. Structural response graph of single degree of freedom

In this section, the dynamic response of a structure of a

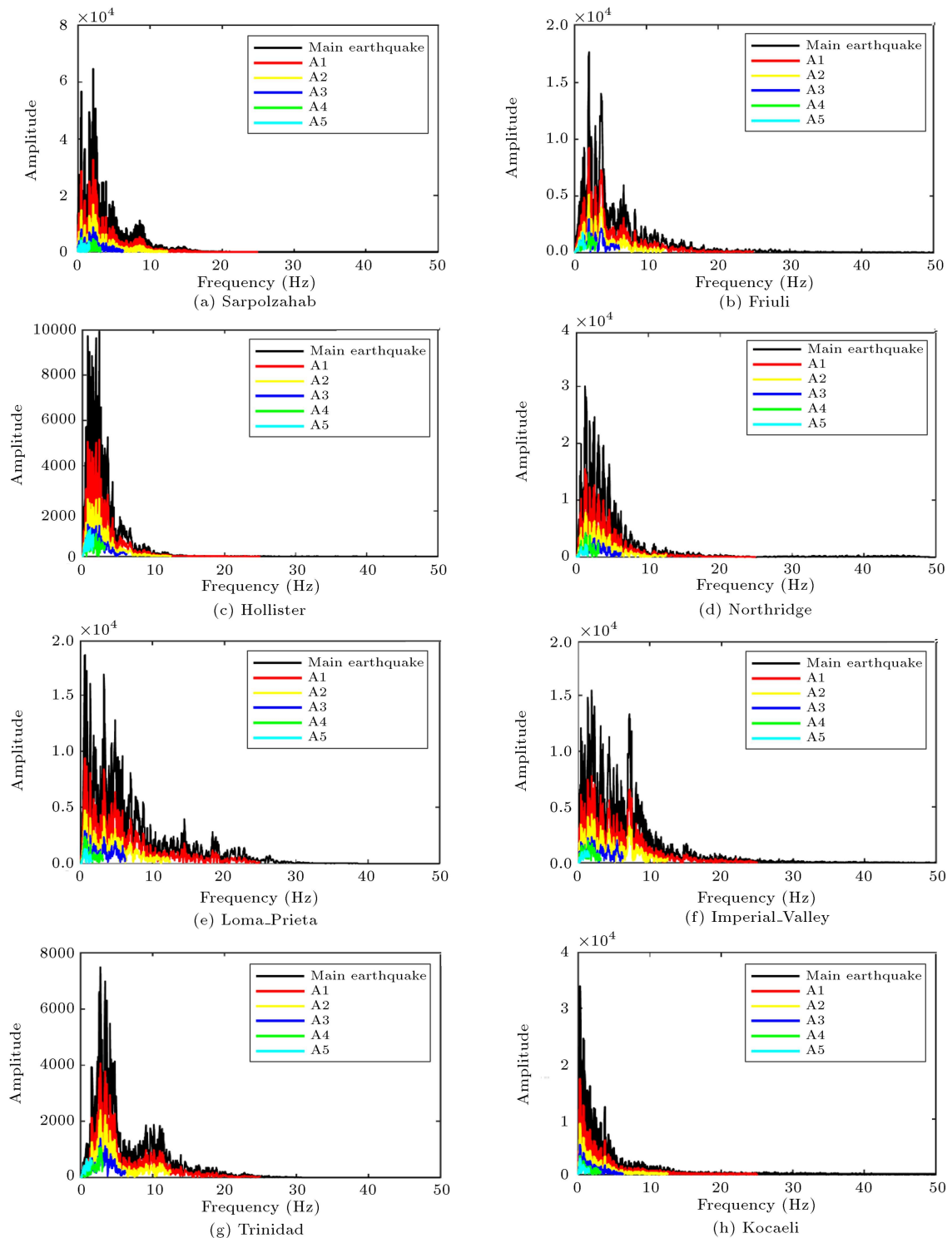


Figure 3. Fourier spectrum of various earthquakes and wavelet filters.

degree of freedom with a damping of 5% and a natural frequency of 2 seconds is checked. The purpose of this is to demonstrate the applicability of the proposed method in time history dynamical analysis. Figure 7 shows the dynamic response of the ideal structure for various waves. Figure 7 shows that the A4 wave is the most effective filtering stage in all earthquakes. This

wave represents a 96% reduction in calculations that are nearly identical to the original earthquake.

9. The earthquake displacement spectrum chart

The displacement spectrum of a structure with a single

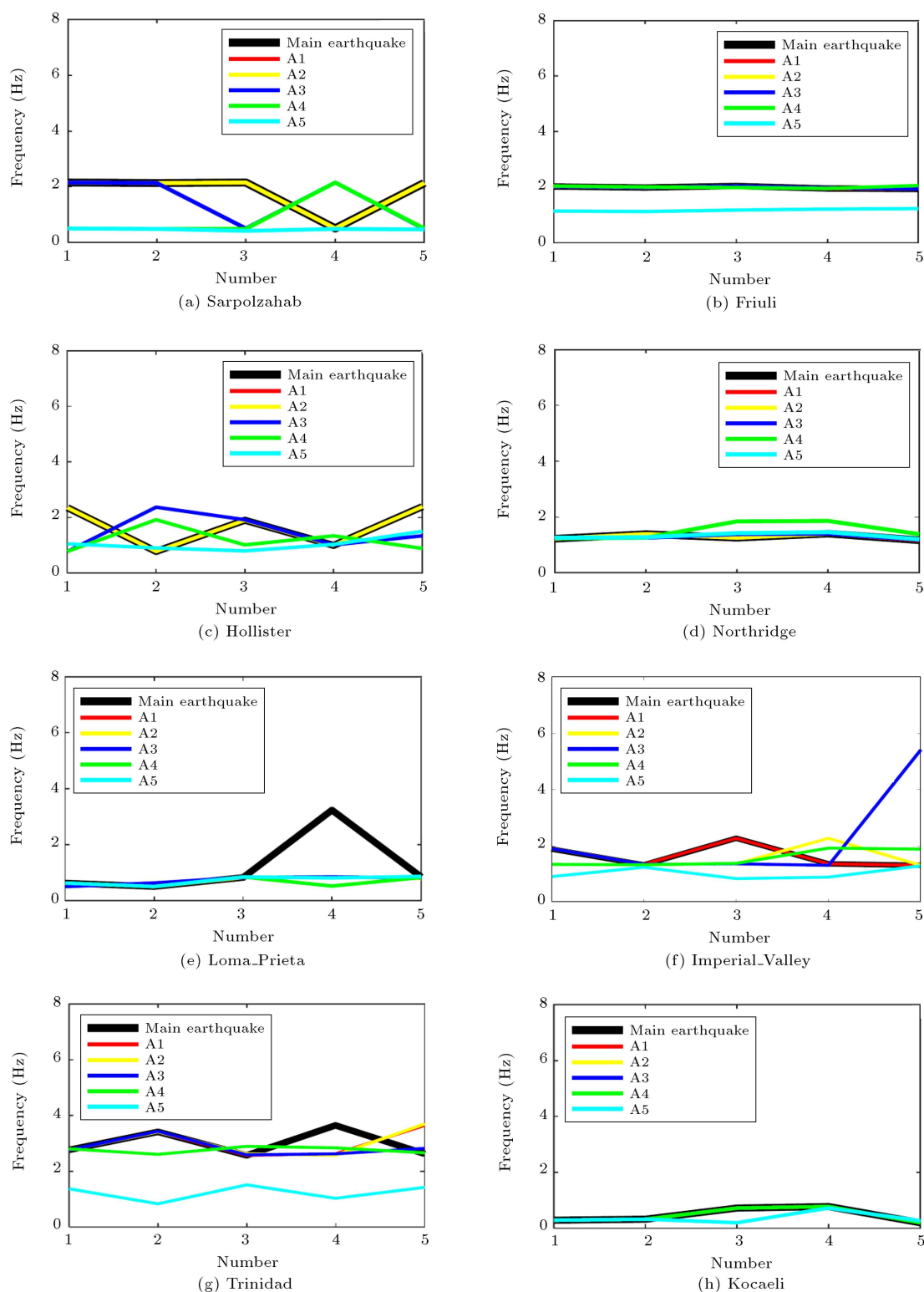


Figure 4. Five dominant frequencies of various earthquakes.

degree of freedom and a 5% damping is examined in this section. As shown in Figure 8, the A2 wave in all earthquakes is a good alternative to a major earthquake, with a 75% reduction in calculations.

10. Conclusion

In this paper, for the first time, the wavelet transform method was used to reduce computations based on seven different concepts. The concepts were chosen in

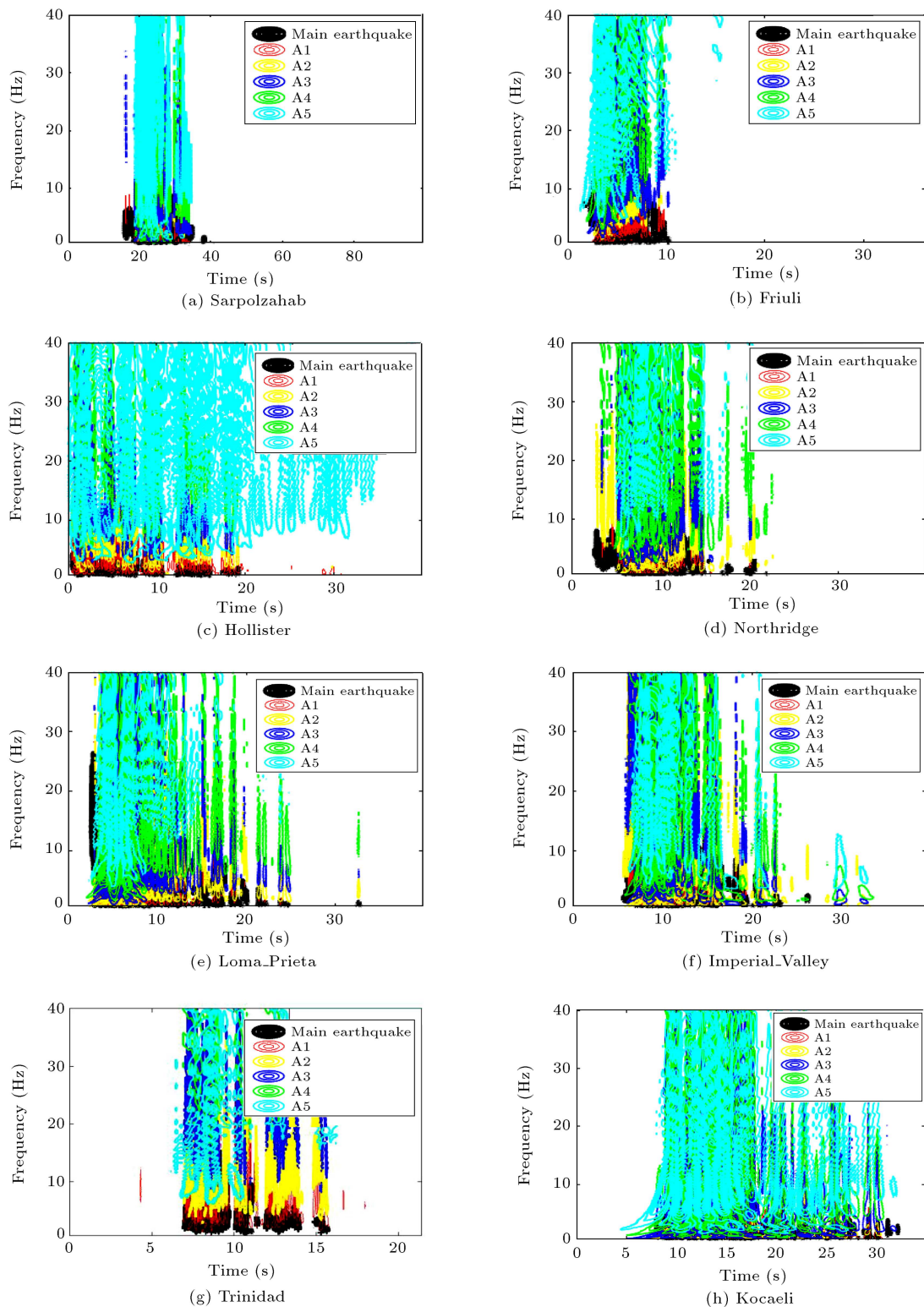


Figure 5. Time-frequency diagram of various earthquakes.

such a way that the frequency content of the waves would affect the results. The aim of this study was to investigate the applicability of the wavelet method. Wavelet transform has a good performance in terms of reducing computations, according to the results.

To this end, the discrete wavelet filter was applied to earthquakes up to 5 stages. The following conclusions can be drawn from the earthquake study:

1. The wavelet filter can be said to be efficient to stage

Table 2. Significant duration of earthquakes.

Desired earthquake	The wave associated with every earthquake	Significant time (sec)		
		Start	End	Total time
Sarpolzahab	Main wave	19.58	30.41	10.83
	A1	19.58	30.56	10.98
	A2	19.58	30.6	11.02
	A3	19.6	30.72	11.12
	A4	19.68	31.2	11.52
	A5	19.52	31.68	12.16
Friuli	Main wave	3.49	7.72	4.23
	A1	3.46	7.92	4.46
	A2	3.48	7.76	4.28
	A3	3.44	7.76	4.32
	A4	3.36	8.16	4.8
	A5	3.20	8.96	5.76
Hollister	Main wave	0.99	15.57	14.58
	A1	0.98	15.56	14.58
	A2	0.96	18.00	17.04
	A3	0.96	17.44	16.48
	A4	0.96	18.08	17.12
	A5	1.92	24.96	23.04
Northridge	Main wave	5.37	14.45	9.08
	A1	5.38	14.44	9.06
	A2	5.40	14.44	9.04
	A3	5.60	14.4	8.8
	A4	5.60	16.64	11.04
	A5	6.40	19.84	13.44
Loma_Prieta	Main wave	4.00	15.31	11.31
	A1	4.10	15.46	11.36
	A2	4.16	15.64	11.48
	A3	4.16	15.92	11.76
	A4	4.16	18.56	14.4
	A5	3.84	16.96	13.12
Imperial_Valley	Main wave	7.08	16.00	8.92
	A1	7.08	16.00	8.92
	A2	7.28	16.20	8.92
	A3	7.76	17.04	9.28
	A4	7.68	20.00	12.32
	A5	8.80	20.48	11.68
Trinidad	Main wave	7.13	14.93	7.80
	A1	7.12	14.94	7.82
	A2	7.12	14.92	7.80
	A3	7.12	15.12	8.00
	A4	7.20	15.04	7.84
	A5	7.36	15.36	8.00
Kocaeli	Main wave	9.83	25.45	15.62
	A1	9.82	25.46	15.64
	A2	9.84	25.48	15.64
	A3	9.84	25.68	15.84
	A4	9.76	25.60	15.84
	A5	9.60	25.60	16.00

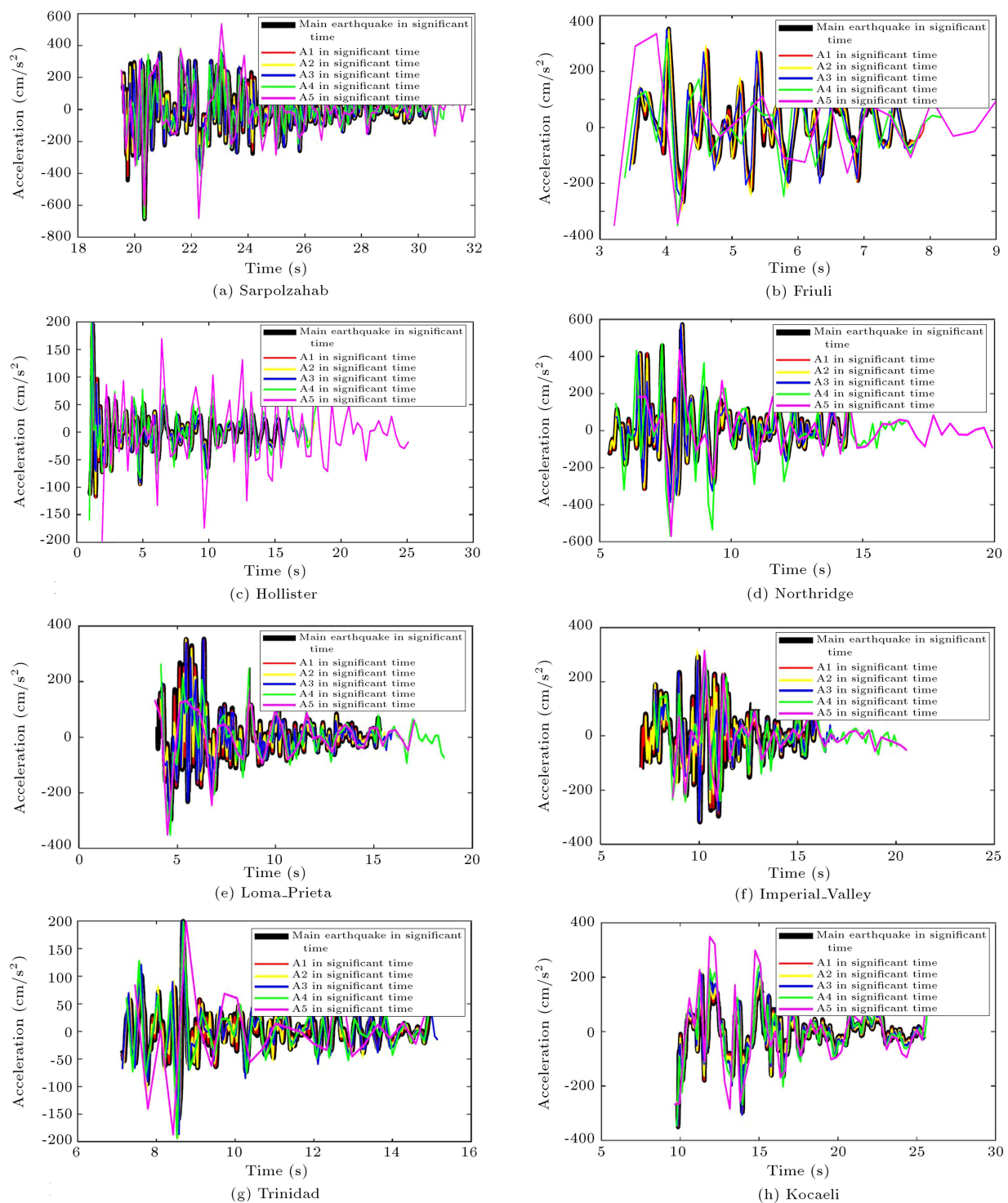


Figure 6. Waves of strong ground motion of various earthquakes.

- 4 based on the acceleration mapping of various earthquakes (A4). Therefore, instead of using the entire earthquake record data, one of their sixteen can be used;
2. The Fourier spectrum of the third level of the wavelet decomposition is close to the main earthquake wave in terms of frequency content;
3. Given the five different earthquake frequencies, it can be seen that an A2 wave is a good option for earthquakes and the calculations are reduced by 75%;
4. The time of the dominant frequencies of all stages of discrete wavelet filters is equal to the time of the dominant frequencies of the earthquake, according

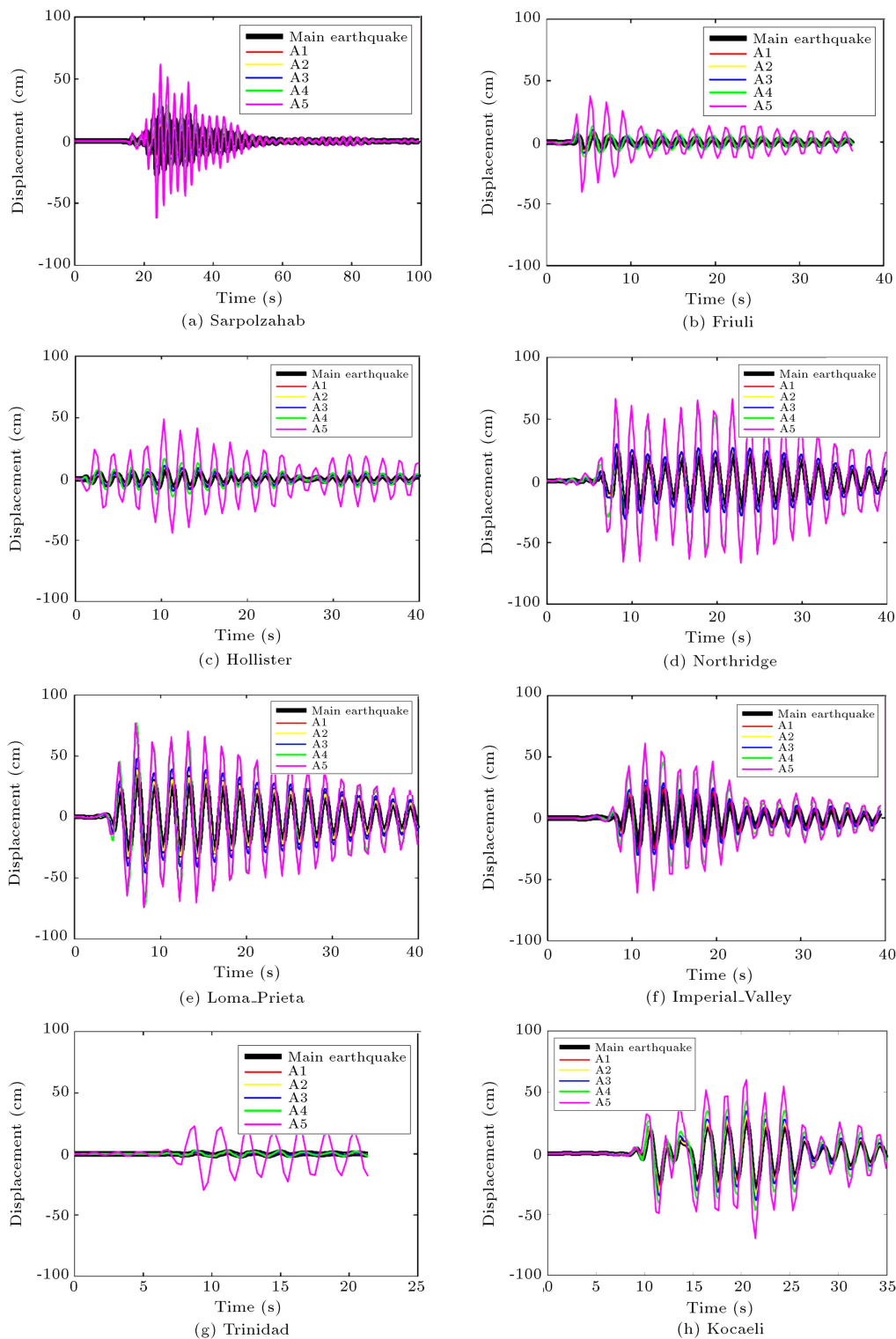


Figure 7. Structural response under the impact of various earthquakes.

to the results derived from the time-frequency curve of continuous wavelet transformation for earthquakes. The wavelet filters and seismic energy are interchangeable;

5. According to the part of the wave that corresponds

to the time of the earth strong motion, the A4 wave curve corresponds to the original earthquake with a reduction of 97% in the calculation;

6. The response curve of a structure of a single degree of freedom indicates that the response curve

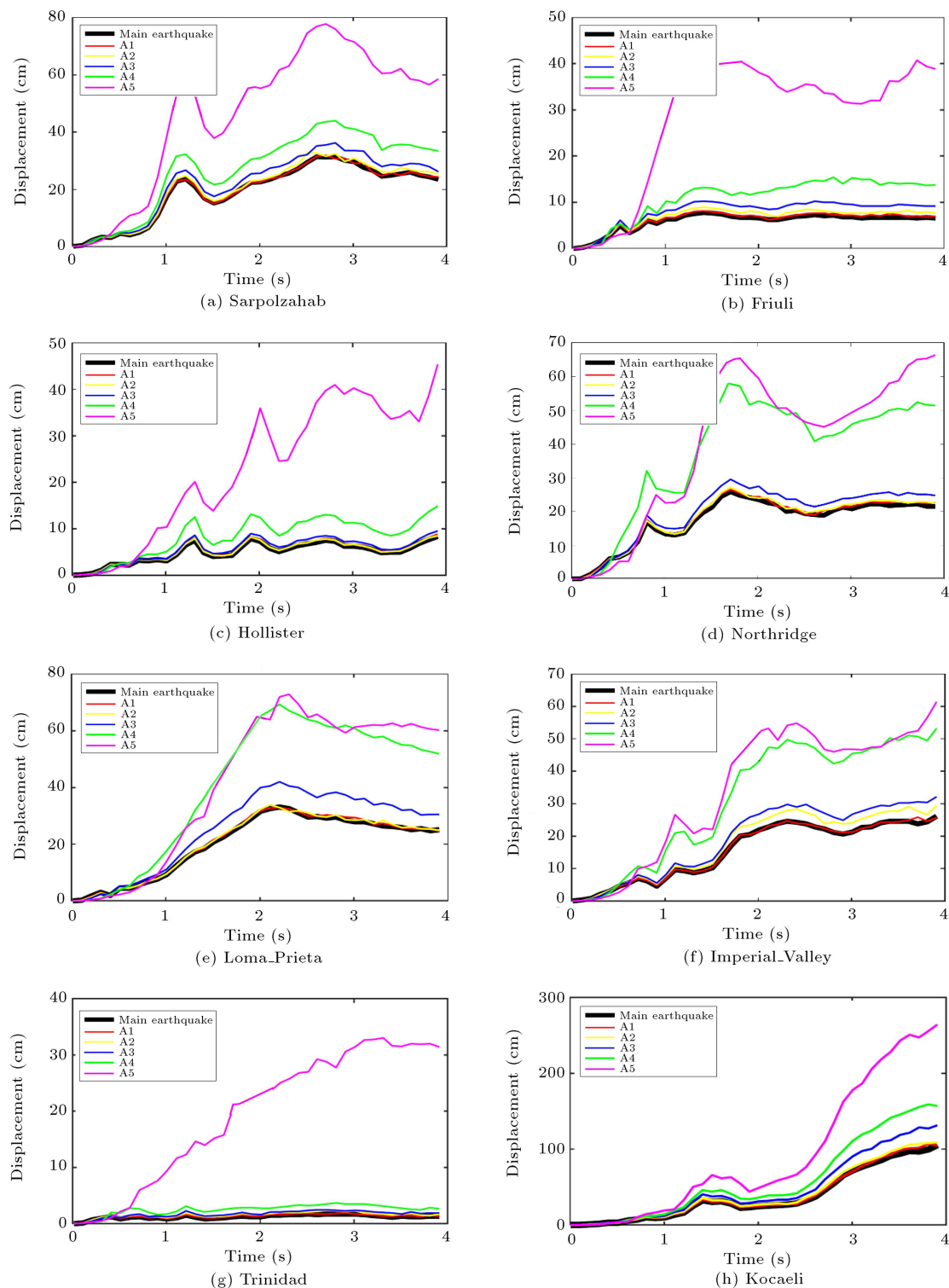


Figure 8. Displacement spectrum of various earthquakes.

obtained from the A4 wave with a reduction of 97% of the calculations is well suited to the response curve from the original earthquake;

7. The curve of the displacement spectrum of a structure of a single degree of freedom indicates that

the response curve obtained from wave A2 with a decrease of 75% of the calculations is well suited to the original earthquake response curve.

In general, the A2 wave, in all of its seismic dimensions, both in terms of frequency content and energy, is

a good substitute for the main earthquake, allowing calculations to be reduced by up to 75%.

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